CHAPTER 127

WAVE EFFECTS ON BUOYANT PLUKES by Nobuo Shuto Chuo University, Tokyo, Japan and Le Huu Ti Asian Institute of Technology Bangkok, Thailand

Abstract

When a buoyant plume is discharged into water where wave motions exist, the axis of plume is bent over and horizontally fluctuated. The dilution rate of the plume is affected by this motion. Experiments are carried out to investigate the dilution rate of plumes when they An analogy to the case of arrive at the free surface. buoyant plumes discharged into uniform steady cross streams is applied to analyse the experimental data. Within the experimental range, it is found that the dilution rate is inversely proportional to the square of the ratio of the water depth to the diameter of outlet and is proportional to the ratio of the discharge velocity to a characteristic horizontal velocity of the ambient fluid. The entrainment coefficient is given as a function of a densimetric Froude number for small Froude number and is a constant, 0.28, for bigger Froude number than 4.3.

Introduction

It is well known that the buoyant plume discharged from

a point source into a cross stream shows a different dilution rate compared with that into still water. The difference is considered due mainly to two causes. The first cause is that the total length of the plume is longer in a cross stream than in still water. Consequently, the plume in a cross stream has a bigger chance to be diluted by the clean ambient fluid. The second reason is that the section of the plume in a cross stream can not be approximated by a circle, but by a "horse shoe" shape which is a result of the bigger upward motion near the central part of the section. In the plume of this kind, the entrainment of ambient fluid is caused not only by the turbulence due to plume motion but also by this upwards motion, while in case of a plume in still water the turbulence is only one major cause of entrainment.

Several researchers have treated buoyant plumes in uniform steady cross streams. The present paper reports experimental results of buoyant plumes discharged into water waves.

Theoretical consideration

It is assumed that the axis of the plume is nearly horizontal, and that the horizontal movement of the plume is almost equal to that of the ambient fluid. The vertical movement of the plume is, however, governed by the buoyancy due to the density difference and the initial momentum, and is quite different from that of the ambient fluid. The difference in the vertical velocities is the main cause to yield the turbulence and entrain the ambient fluid. The entrainment is expressed by a coefficient $\overline{\alpha}$ which is the ratio of the entrainment velocity to the velocity difference.

The present paper does not treat a detailed structure of the plume, but only discusses values on the axis of the plume, which are considered important from the practical point of view.

2200

Sections of the plume are assumed circle. The velocity and density distributions are assumed to take the "top-hat" shape and follow a similarity law. Caution is needed to the difference bet ween these assumptions and the reality. Although we have just mentioned that the upward motion of water near the center of the plume was one of the main causes of bigger entrainment, no consideration of the phenomenon is taken into the above assumptions. All the effects are included in values of α which is to be experimentally determined.

The following four equations, integrated over a section of the plume, are basic equations which still contain the unsteady terms.

Equation of continuity

$$\frac{\partial D^2}{\partial t} + U \frac{\partial D^2}{\partial x} = 4\alpha Dw \tag{1}$$

Equation of motion in vertical direction

$$\frac{\partial (D^2 w)}{\partial t} + U \frac{\partial (D^2 w)}{\partial x} = \frac{\Delta \rho}{\rho} g D^2$$
(2)

Equation of conservation of density defficiency $\partial (D^2 \Delta \rho) = D^2 \Delta \rho$

$$\frac{\partial t}{\partial t} + U \frac{\partial t}{\partial x} = 0$$
(3)

Path of plume

$$\frac{\partial y}{\partial x} = \frac{w}{U} \tag{4}$$

in which α denotes the entrainment coefficient, D the diameter of the section of plume, w the vertical velocity in the plume, U the horizontal velocity of the ambient fluid, ρ the density and $\Delta\rho$ the density difference.

Since it is not easy to solve these equations analytically, the following consideration is introduced to obtain approximate expressions.

The plume discharged while the ambient fluid is flowing rightwards will rise diagonally rightwards. Then, the direction of motion of the ambient fluid changes leftwards. The plume, too, changes the direction of its ascent. New part of the plume which now begins to be discharged locates

1...

below the old part of the plume, and these two parts of the plume do not interact each other. If the old part of the plume is reflected with respect to a vertical line at which the plume arrives at the end of the rightward movement, the whole plume is regarded similar to the plume in a unidirectional stream. Accordingly, if a characteristic velocity of the ambient fluid is well defined, the theoretical considerations used in the analysis of plumes in a cross stream can be applied with a sufficient accuracy to the present case in order to estimate the dilution rate of the plume at the free surface.

We select the mean velocity U averaged over a half wave period as a representative velocity. Within the range that the wave motion is approximated by long waves, the horizontal velocity is vertically uniform.

These assumptions are used in averaging the above set of equations. Such terms as D and w which vary periodically with respect to time may have some contributions through the interaction terms such as D^2w to the analysis. For example, the initial velocity, w_0 , of the plume can vary The discharge of the plume is determined by periodically. the piezometric head between the constant head tank and the outlet of the plume at the bottom of channel. Water pressure in the channel fluctuates to some extent due to wave motion, and the term D^2w does not vanish after the The fluctuation is, however, considered small averaging. enough to be neglected in the analysis. Thus, contributions of such terms are included in the value α , which can be a function of F_{rv} or F_{rw_n} due to this approximation.

Averaging the equations (1) to (4), we have the same set of equations as in the case of the plume in a uniform cross stream, and the solutions are given in the reference (4).

Dilution rate at the free surface is given by

$$\frac{\Delta\rho}{\Delta\rho_0} = \frac{1}{4a^2} k \left(\frac{h}{D_0}\right)^{-2}$$
(5)

where h denotes the water depth and D_0 the diameter of the outlet.

Experimental equipment and procedure

A channel, 7.4 m long, 75 cm high and 50 cm wide, at the Asian Institute of Technology was used in the experiment. Two piston-type wave generators placed at each end of the channel can be operated synchronously to generate standing waves. The channel was filled with salt water.

Fresh water is storaged in the two tanks, each with a capacity of 180 litres and pumped up to the constant head tank, from which fresh water is discharged through a 5/8 in. pipe. Water overflowed the constant head tank returns to the storage tank and is repumped up.

The outlet is placed at the point where we have the maximum horizontal velocity under standing wave motion. Diameters of the outlet used in the experiments are 0.5 cm, 0.75 cm and 1.0 cm.

The discharge is determined by measuring the decrease of water quantity in the storage tank.

In order to know the dilution rate, samples were taken near the water surface, and their salinity and concentration of chromium were measured. Initial concentration of chromium of the fresh water was $5000 \ \mu g/l$.

A spectrophotometer was used to determine the concentration of chromium of the samples, and as the accuracy of this method was better than that of the measurement of salinity, only the results obtained by the spectrophotoanalysis are used in the following analysis.

The samples are taken from three points; one point just above the outlet, two points 50 cm apart from this point up- and downstream along wave direction. After having examined that the experimental condition arrived at steady condition and having stopped the wave generator, water samples were taken. Thickness of the surface layer which was composed of the diluted plume were 3 to 5 cm, and the inlets of the samplers were inserted at 1-2 cm below the free surface. Intake velocity was controlled not to entrain the water from the lower layer. No sampling was carried out at deeper points because the sampling only from the plume itself is impossible due to the horizontal fluctuation of the path of the plume caused by wave motion.

Wave periods were determined from the rotation of the wave generator, and wave height was measured from its spacial distribution recorded along the side wall.

A mixture whose density is the same as that of water was made from benzene and carbon tetrachloride, and small particles of the mixture were introduced in the tank to measure the maximum excursion distance. Theoretical estimate of the excursion distance given by the theory of long waves of small amplitude agreed very well with the experimental results.

Although standing waves were used in the experiment, it is possible to transfer the present results to that under progressive waves, because the excursion length is so small compared with the wave length that the spacial variations of horizontal velocity do not affect so much the dilution of the plumes.

Experimental conditions were as follows.

$$h = 12.5, 25, 35 \text{ cm}$$

$$h / D_0 = 12.5 \sim 70$$

$$F_{rv} = 1 \sim 7$$

$$k = \frac{w_0}{U} = 3 \sim 30$$

$$w_0 = 0.32 \sim 3.00 \text{ m/sec}$$

$$\Delta \rho_0 = 0.01 \sim 0.05 \text{ g/cc}.$$

$$H = 1.4 \sim 5 \text{ cm}$$

$$T = 2.9 \sim 6.0 \text{ sec}.$$

where F_{rv} is a densimetric Froude number defined with the horizontal velocity U, H the wave height of progressive waves and T the wave period.

2204

Experimental results and discussions

(1). Classification of plumes.

Plumes are classified according to how many times the wave motion is repeated until they arrive at the free surface. A plume is called "nearly vertical" if the plume reaches the free surface within one half wave period. An "intermediate" plume is one that reaches the free surface within one wave period. Others are called "nearly horizontal". Thickness of the surface layer produced by the "nearly vertical" plume was the smallest, and the "nearly horizontal" plume gave the thickest surface layer. Figure 1 shows the classification of the plume in terms of the initial velocity of the plume, the horizontal velocity of the ambient fluid, the diameter of the outlet and the water depth. Except for the cases of smaller relative depth than 30, major factor which determines the type of plume is the relative velocity.

(2). Dilution rate of the plume at the free surface.

Abraham gave a formula of dilution for plume in still water as follows;

$$\frac{\Delta \rho}{\Delta \rho_0} = 9.7 F_{rw_0}^{2/3} \left(\frac{h}{D_0} + 2\right)^{-5/3} \tag{6}$$

which might be a good approximation for the case that the plume rises nearly vertically under small or negligible effects of the horizontal velocity of the ambient fluid. If the phenomenon is well approximated by the plume in a cross stream, the ratio $\frac{d\rho}{d\rho_0}$ may be proportional to $(h/D_0)^{-2}$.

The least scatter of experimental results obtained is shown in Fig.2, where the empirical relationship is

$$\frac{\Delta \rho}{\Delta \rho_0} = 0.17 F_{rw_0}^2 \, k^{-1} \left(\frac{h}{D_0}\right)^{-2} \tag{7}$$

or

$$\frac{\Delta\rho}{\Delta\rho_0} = 0.17 F_{r_U}^2 k \left(\frac{h}{D_0}\right)^{-2} \tag{8}$$

Comparison with Eq.(5) gives

$$\alpha = 1.21 F_{r_U}^{-1}$$
(9)

In the above equations, F_{rv} and F_{rw_0} are the densimetric Froude number, with respect to the horizontal velocity of the ambient fluid U, and the initial velocity of the plume w_0 , respectively.

(3). Entrainment coefficient.

Figure 3 shows α as a function of F_{rv} , after evaluating it by Eq.(5). For large values of F_{rv} , α is better approximated by a constant, 0.28, which is a little smaller than 0.33 for cases of the plume in a cross stream.

The major reason of this difference seems to depend on the difference of the definition of $U_{\: \bullet}$

For smaller values of F_{rv} , α is expressed as a function of F_{rv} . Fan gave

$$\alpha = f(k F_{rw_0}^{5/8}) \tag{10}$$

but in our cases, it is discovered more appropriate to express α as function of F_{rv} .

Cares are needed when we use the results given in Fig. 3. For α is bigger than unity, for very small values of $F_{\tau \nu}$. The coefficient α is the ratio of the entrainment velocity to the velocity which excites the entrainment. It is, therefore, physically impossible that α can exceed unity without any amplification mechanism. This fact tells that our model fails.

In addition, caution should be paid on the fact that values of α scatter very much for $h/D_0 < 16$. In the vicinity of the outlet, no similarity law is applicable due to the effect of high initial velocity. This affects the values of α , too.

Even with these faults, results shown in Fig.3 are practically applicable to compute the dilution rate of the plume.

2206

(4). Empirical formula of the dilution of the plume.

Within the present experimental range where wave motion is sufficiently approximated by long waves of small amplitude, U in Eq.(8) is rewritten in terms of h and H and we have

$$\frac{d\rho}{d\rho_0} = \frac{1}{1.27} \frac{1}{\alpha^2} \frac{w_0}{\sqrt{gh}} \frac{h}{H} \left(\frac{h}{D_0}\right)^{-2}$$

$$\alpha = 1.21 F_{rU}^{-1} \quad F_{rU} \leq 4.3$$

$$= 0.28 \quad F_{rU} \geq 4.3$$
(11)

Conclusions

An empirical formula, Eq.(11), is established to compute the dilution rate of the plume discharged into wave motion which can be approximated by long waves.

Entrainment coefficient, α , is also given by Eq.(11), but caution is needed because the physical basis of α for small Froude number is ambiguous.

Nevertheless $\mathbb{E}q_{\bullet}(11)$ is sufficient enough for practical purposes.

References

- 1. Abraham, G. (1960): Jet diffusion in liquid of greater density, J. Hyd., Div., Proc. ASCE, Vol. 86, HY 6, pp.1-13.
- Fan,L.N. (1967): Turbulent buoyant jets into stratified or flowing ambient fluids, Rep. KH-R-15, W. Keck Lab., California Institute of Technology.
- Morton, B.R. et al. (1956): Turbulent gravitational convection from maintained and instantaneous sources, Proc. Roy. Soc. London, Ser.A, Vol.234, pp.1-23.
- 4. Shuto, N. (1971): Buoyant plume in a cross stream, Coastal Engineering in Japan, Vol.14, pp.167-173.
- 5. Slawson, P.R. and G.T. Csanady (1967): On the mean path of buoyant, bent-over chimney plumes, J. Fluid Mech., Vol. 28, Part 2, pp.311-322.



Fig.1 Classification of plumes

Fig.3 Entrainment coefficient



Fig.2 Dilution rate of plumes at free surface